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Research Paper

Geoelectric Evaluation of Groundwater Potential and Aquifer Vulnerability of Overburden Aquifers of Ado-Ekiti, Southwestern Nigeria.

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ABSTRACT: Groundwater resources are crucial for the socio-economic development of regions with seasonal surface water sources. Characterization of groundwater aquifers typically involves assessing groundwater potential and vulnerability. This study utilized 40 Vertical Electrical Sounding (VES) measurements with a Schlumberger array to characterize the basement aquifers of Ado-Ekiti, Southwest Nigeria. The interpretation of the data was effected using IPi2Win and Interpex (IX1D). The findings reveal four geoelectric layers in 67.5% of the study area and five layers in 32.5%. The five identified layers include topsoil (0.24-21m; 0.97-327.76 Ωm), clay/laterite (0.89-4.43m; 6.22-368.07 Ωm), weathered basement (0.32-117.3m; 9.73-5705 Ωm), fractured basement (0.58-46.94m; 1.64-2200.2 Ωm), and fresh unweathered basement (23.47-48502 Ωm). Aquifer layers, consisting of weathered and fractured basement, have thicknesses ranging from 0.9 to 134.4m, while the protective layers of topsoil and clay range from 0.2 to 21.1m. 52.5% of the study area exhibits good groundwater potential, with hydraulic conductivity and transmissivity ranging from 0.3 to 53.7m² /day and 41.7 to 4280.4m²/day, respectively. Groundwater potential in the study area is classified based on the depth to the basement, with 17.5% of the area rated as negligible (depth < 10 m), 30% as low (10–20 m), 32.5% as moderate (20–30 m), and 20% as good (depth > 30 m). This indicates that 52.5% of the study area exhibits moderate to good groundwater potential, even when considering saprolite and fracture characteristics. Geoelectric crosssections identify areas of high groundwater potential in valleys filled with weathered material. Longitudinal conductance ranges from 0.07 to 8.99 mho, with 70% of the area classified as poorly to very poorly protected, suggesting that localized management strategies may be necessary to prevent aquifer contamination. KEYWORDS: Aquifer, Vulnerability, Longitudinal Conductance, Transmissivity, groundwater

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I. INRODUCTION

Groundwater serves as an important resource for domestic, industrial and agricultural uses, especially in regions where surface water is scarce and unreliable. In the basement complex terrains, the evaluation of groundwater potential is challenging due to the heterogeneous and anisotropic nature of the underlying rock formations. Assessing groundwater potential in basement complex regions requires an integrated approach that considers the geological, geophysical and hydrological properties of the area.

Groundwater occurrence in Ado-Ekiti is majorly from fissure, joints, weathered and fractured zone of crystalline rocks of older granite, charnokites or metamorphic origin of Precambrian age. The void and joints are interconnected and this enables the movement of the groundwater and consequently the passage of contaminants from the overburden [1]. Generally, the movement of groundwater is influenced by topographical elevation and pressure.

Contaminants from the surface through run off can easily leak into aquifer depending on the bearing capacity of the soil. Weathered loose materials over lay aquifer expose ground water to contamination while aquifers overlain and underlain by confining layers are less vulnerable to contamination. Aquifer is generally overlain by different soil material of variable thickness, the soil material include lateritic soil, clay, sandy and silt –sized particles deposited by run off and surface water [2]. These aquifers are localized and of low porosity and permeability, the occurrence also depends on secondary porosity.

II. THE STUDY AREA AND HYDRO-GEOLOGICAL FORMATION

The area of study covers Ado-Ekiti municipal, Ekiti state Nigeria. It is within 733000mE to 765000mE and between 835000mN to 855000mN covering approximately $640km^2$. Ado Ekiti is situated at the centre of Ekiti region of western Nigeria. It is bounded in the South and East by Ondo state through Ikere Ekiti Local Government Area, in the North by Kwara state and Kogi state, in the West by Osun state. The population is about 427,700 according to 2016 population censor; the elevation is between 334m and 510m.

The geology of Ado-Ekiti area is dominated by crystalline and schistose rocks which belong to the Southwestern Nigeria basement complex [3]. The lithology in Ado-Ekiti area is made up of migmatite-gneiss, granite gneiss, charnockite, older granite and metasediments which are quartzite-muscovite schists. Field study shows that the charnockite and granite are closely associated in time and space [4]. According to [5] both granites and charnockites are contemporaneous or the charnockites formed shortly after the emplacement of the older granites. Older metasediments occur as ridges and mica schists and this constitutes part of the Schist belts in Nigeria [3] [6]. Ado Ekiti is surrounded by hills and inselbergs of different shapes in ridges, a rugged topography with rivers and streams meandering through the valleys. The major Rivers are Ireje, Omisanjana and Elemi that flow to River Ose and Owena in Ondo State at the southern part of Ekiti, this drain down to Atlantic Ocean in the southern part of Nigeria [7].

Figure 1: Geological map of Ado Ekiti Southwestern Nigeria [4]

III. MATERIALS AND METHODS

3.1 Vertical Electrical Sounding (VES)

This study utilizes the Electrical Resistivity (ER) method, employing the Schlumberger array setup to perform Vertical Electrical Sounding (VES). ER surveys are based on the principle that water enhances the conductivity of rocks, thereby reducing their resistivity [8]. The Schlumberger array configuration consists of two inner potential electrodes (MN) and two outer current electrodes (AB). The spacing between the potential electrodes is relatively small compared to the spacing between the current electrodes, which allows for the exploration of deeper subsurface resistivities as the current electrodes are spread farther apart. During the survey, the distance between the current electrodes (AB) is gradually increased. Although the potential electrodes may be slightly adjusted to maintain adequate signal strength, their spacing (MN) remains mostly unchanged. The

guideline often followed is that AB should be at least five times MN [9] and [10]. The Schlumberger array configuration is particularly beneficial because it reduces the effects of near-surface heterogeneities, offering a more accurate representation of deeper subsurface resistivity variations.

The inversion of the vertical electrical resistivity (VES) data was carried out using Ipi2Win for data smoothing and then Interpex IX1D v3, which accurately computes layer boundaries and resistivities through an iterative curve-matching approach. This method involves iteratively adjusting model parameters, such as layer thickness and resistivity, to reduce the difference between the observed field data and the theoretical model curve.

3.2 Secondary Aquifer and Hydraulic Parameters

The secondary parameters calculated to interpret the inverted VES data in terms of aquifer potential and vulnerability include Dar Zarrouk parameters (longitudinal conductance and transverse resistance), the coefficient of anisotropy (λ) , hydraulic conductivity (K) , and transmissivity (T) . These parameters were computed using the:

 ρ is aquifer resistivity

3.3 Groundwater Potential

Groundwater classification based on transmissivity, as outlined by [14], categorizes aquifers into three potential classes: low $(T < 50 \text{ m}^2/\text{day})$, moderate $(T = 50-500 \text{ m}^2/\text{day})$, and high $(T > 500 \text{ m}^2/\text{day})$ (Table 1). The groundwater potential was assessed using the depth to bedrock function proposed [15], where depths of $\langle 10 \text{ m} \rangle$ 10-20 m, 20-30 m, and >30 m were rated as negligible, low, moderate, and good potential, respectively (Table 2). The potential of the weathered aquifer was evaluated using the saprolite resistivity function proposed by [16]. In this method, resistivity values of ≤ 20 Qm were rated poorly due to potentially high clay content, while values between 20-100 Ωm indicated optimal groundwater potential. Resistivity values of 100-150 Ωm, 150-300 Ωm, and >300 Ω m were rated as medium, limited or poor, and negligible potential, respectively (Table 3). For the fractured bedrock aquifer potential, based on the function from [15], resistivity values of <750 Ωm, 750-1500 $Ωm$, 1500-3000 $Ωm$, and >3000 $Ωm$ were rated as high, medium, low, and negligible potential, respectively (Table 4).

 $T = K \cdot h$ (4)

3.4 Protective Capacity

Longitudinal conductance is employed to assess the protective capacity of the overlying layer, following the classification scheme proposed by [17]. In this classification, values greater than 0.5 mho suggest high protection, values between 0.5 to 0.1 mho indicate fair protection, values from 0.1 to 0.05 mho reflect poor protection, and values less than 0.05 mho denote very poor protection (Table 3.5).

Maps displaying isoconcentrations of various parameters were created using Golden Software Surfer 18, with computations performed in Microsoft Excel 2010. Geoelectric cross-sections were developed using Adobe Illustrator.

Table 5: Aquifer Protective capacity classification based on Longitudinal Conductance (mho) values [17]

Longitudinal Conductance (mho)	Protective Capacity Classification		
>0.5	High Protection		
$0.1 - 0.5$	Fair Protection		
$0.05 - 0.1$	Poor Protection		
< 0.05	Very Poor Protection		

IV. RESULTS AND DISCUSSION

4.1 Resistivity Data

The study identified between four and five geoelectric layers across the area, with the majority of sites (67.5%) showing four layers and the remaining 32.5% showing five layers. The identified layers include topsoil, clay/laterite, weathered basement, fractured basement, and fresh unweathered basement. This interpretation was supported by observations from hand-dug wells examined during the survey.

Nine distinct curve types were identified in the study. For the four-layer sites, the curve types include AA (5%), HA (22.5%), KH (20%), and QH (20%). In the five-layer sites, the curve types are AKQ (5%), HAA (5%), HAK (2.5%), HKH (17.5%), and KHK (2.5%). In basement terrains, lower resistivity values in deeper geoelectric layers—represented by curve types with "K" and "Q" in their nomenclature—are typically associated with higher groundwater potential. This suggests that 67.5% of the area, characterized by curve types such as KH, QH, AKQ, HAK, HKH, and KHK, could potentially be viable locations for high groundwater yield, provided the thickness of the low-resistivity layer is sufficient.

The topsoil layer shows a thickness range of 0.24 to 21 m and a resistivity range of 0.97 to 327.76 Ω m. This wide range in resistivity is primarily attributed to varying soil moisture content, with lower resistivity values indicating the presence of clayey and moist soil, while higher resistivity values suggest sandy or dry lateritic conditions. The clay/laterite layer serves as a protective barrier for the underlying aquifers, with thicknesses ranging from 0.89 to 4.43 m and resistivity values between 6.22 and 368.07 Ωm. This layer is vital for reducing contamination risks by preventing surface pollutants from infiltrating the groundwater. The weathered basement, with a thickness ranging from 0.32 to 117.3 m and resistivity values between 9.73 and 5705 Ω m, is a crucial component of the aquiferous layer in the region. Weathered basement zones with lower resistivity typically suggest higher porosity and permeability, which are favorable for groundwater storage and transmission. Conversely, areas with higher resistivity are indicative of more competent, less permeable materials. These findings align with several studies in the region. The fractured basement, with thicknesses ranging from 0.58 to 46.94 m and resistivity values between 1.64 and 2200.2 Ω m, plays a vital role in groundwater storage within the study area. This finding is consistent with previous research, where the fractured basement has been recognized as a key aquifer unit in basement terrains.

Figure 2: Vertical electric sounding curve and layers (KH-Curve Type, VES 23)

Figure 3: Vertical electric sounding curve and layers (HKH-Curve Type, VES 8)

4.2 Groundwater Potential

The depth to the top of the basement varies significantly, ranging from 1.16 m at VES 13 to 138.73 m at VES 35. Based on the depth-to-basement function proposed by [15], the groundwater potential across the VES positions is categorized as follows: 17.5% are rated as negligible $(d < 10 \text{ m})$, 30% as low (10-20 m), 32.5% as moderate (20-30 m), and 20% as good ($d > 30$ m). This implies that a substantial portion (21 out of 40 VES) of the area exhibits moderate to good groundwater potential, particularly where the depth to the basement exceeds 20 meters. However, the presence of zones with negligible and low potential suggests that groundwater availability may be uneven. The weathered layer, where the depth to the basement exceeds 20 meters, was further categorized using [16] saprolite resistivity function. The results reveal that 52.38% of the areas are rated as optimum (20 < ρ < 100 Ωm), 19.05% as poor (150 < ρ < 300 Ωm), 23.81% as negligible (> 300 Ωm), and 4.76% as poorly rated due to high clay content ($\rho < 20$ Ωm). According to [15], fractured bedrock resistivity function, almost all (97.5%) of the fractured aquifers at these VES points exhibit high groundwater potential (\leq 750 Ω m). This indicates that where the weathered aquifer is not viable, the fractured aquifer could serve as a viable alternative for groundwater development.

Table 10: Aquifer Potential rating based on Saprolite resistivity function

Table 11: Aquifer Potential rating based on Fractured bedrock function

Table 12: Aquifer Potential rating based on Depth-to-basement function

4.3 Protective capacity and vulnerability

The study reveals that the longitudinal conductance of the protective layers (clayey topsoil) ranges from 0.07 to 8.99 mho. 70% of the area is classified as poorly to very poorly protected, making it highly vulnerable to contamination. This may be due to differences in clay content and layer thickness across different regions. In areas with low protective capacity, localized management strategies are essential to prevent contamination from surface pollutants.

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VES 29	ABUAD	KH	0.01	1354.8	Poor
VES 30	Oke-Ila Rd	KH	0.02	592.7	Poor
VES 31	Oke-Ila Rd	KH	0.21	501.8	Moderate
VES 32	Oke-Ila Rd	QH	0.13	673.8	Weak
VES 33	Oke-Ila Rd	QH	0.12	6750.6	Weak
VES 34	Agric Training Centre	KH	0.01	8889.9	Poor
VES 35	Ayegunle	KH	0.29	9868.3	Moderate
VES 36	Olaoluwa Area	QH	0.01	1141.1	Poor
VES 37	Olaoluwa Area	QH	0.01	1037	Poor
VES 38	Olaoluwa Area	QH	0.12	345.6	Weak
VES 39	Igirigiri	HA	0.11	392.6	Weak
VES 40	Igirigiri	KH	0.37	2158	Moderate

Figure 5: Protective Layer Longitudinal Conductance (mho) of the overburden geoelectric layer.

V. CONCLUSION

The study provides a comprehensive assessment of the groundwater potential, and vulnerability of Ado-Ekiti. The aquiferous layers, consisting of the weathered and fractured basement, exhibit moderate to high groundwater potential, particularly in regions with thick overburden and deep basement valleys. The poor protective capacity of the topsoil and clay layers raises concerns about the vulnerability of the aquifers to contamination. Localized management strategies will be essential to protect these valuable groundwater resources, especially in areas with high groundwater potential and poor protection.

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